# Ecology of glacier-fed rivers: current status and concepts

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# SUMMARY

1. This paper is an introduction to a special issue of Freshwater Biology containing papers dealing with various aspects of the ecology of glacier-fed rivers.

2. Using similar field protocols, a wide range of glacier-fed systems were studied across Europe from the French Pyrenees to Svalbard within the framework of the European Commission project, Arctic and Alpine Stream Ecosystem Research (AASER). Recent investigations from other parts of Europe together with New Zealand and Greenland are also reported. This work has advanced our knowledge of the functioning of these types of rivers and has led to the modification and quantification of the conceptual model of Milner & Petts (1994).

3. Glacier-fed rivers, by virtue of the dominance of physical variables in shaping macroinvertebrate communities, are not only good indicators of climate change, but also may be suitable testing grounds for examining ecological concepts.

Between the permanent snowline and the treeline,

three general types of stream ecosystem have been

identified (Steffan, 1971; Ward, 1994): the kryal, or

glacier-melt dominated system; the rhithral, or

seasonal snowmelt-dominated system; and the kre-

nal, or spring-fed system. However, valley glaciers

can extend through the treeline and glacier-melt

dominated streams can persist for considerable

distances downstream reaching sea-level at higher

latitudes (Brittain et al., 2001; Lods-Crozet et al.,

2001a). Furthermore, snow and ice cover may vary

significantly over small spatial scales, and different

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## Introduction

Ecological zonations within cold regions may be defined on the basis of general climatic criteria. The treeline in undisturbed environments approximates the position of the mean 10 °C July isotherm and forms the dividing line between subalpine/subarctic woodlands and the treeless alpine/tundra areas (Remmert, 1980). Above the treeline, the environment is generally characterized by snow beds, bare rock surfaces, skeletal soils and prostrate vegetation except close to river margins, where riparian growth can be extensive especially in the low alpine. Permanent snowfields and glaciers dominate at higher altitudes.

the low alpine. ers dominate at stream segments will display characteristics that reflect the relative proportions of the different runoff sources: glacier melt, snowmelt and spring-flow. Each source generates a characteristic seasonal hydrological signature and runoff of different physical (temperature and suspended sediment load) and chemical quality (Sharp *et al.*, 1998; Brittain *et al.*, 2000).

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#### Concepts

High

From a synthesis of European and Alaskan literature, Milner & Petts (1994) proposed a conceptual model (Fig. 1) to predict the gradient of zoobenthic communities in rivers downstream of glacial margins as determined by two principle variables, water temperature and channel stability. Water temperature conveys broadly predictable trends to the longitudinal pattern of invertebrate community structure in glacial rivers and determines the point downstream at which certain taxa may colonize within a stable channel (Milner, 1994). However, even if water temperature is suitable, channel instability may delay colonization and allow some cold tolerant taxa to remain dominant in the community. Chironomid larvae of the genus Diamesa were suggested as the typical sole inhabitants when  $T_{\text{max}} < 2 \text{ °C}$ , but are joined by other Diamesinae, Orthocladiinae and Simuliidae when  $T_{\text{max}} > 2 \text{ °C} < 4 \text{ °C}$  (Milner & Petts, 1994; Ward, 1994).

This conceptual model also recognizes that, discontinuities in the downstream gradient of zoobenthic communities may be caused by lakes. These act as thermal and hydrological regulators, and by changing stream valley morphology. Lakes embedded in the river course also influence solute loadings, turbidity patterns, and geomorphological dynamics of downstream river sections (see also Ward & Stanford, 1983). Tributaries may act in a similar way. For example, glacial tributaries may retard zoobenthic community development, while non-glacial tributaries may enhance it (Brittain *et al.*, 2001). The general theme outlined by Milner & Petts (1994) is further superimposed by variations in the structure and functioning of stream ecosystems both between and within catchments.

Differences in stream types within arctic and alpine areas typically relate to varying hydrological regimes during the summer, which affect in particular suspended sediment load, turbidity, hydraulic stress and bedload transport. During the winter fewer differences exist as many streams and rivers are dominated by groundwater inputs (Tockner *et al.*, 1997; Malard *et al.*, 1999). Other factors, such as the extent of snow and ice cover, may nevertheless give rise to differences in incident radiation, water temperature and primary production (Schütz *et al.*, 2001). These points highlight the range of stream types found in high latitude and altitude environments, although they fail to link the fundamental hydrological effects of runoff on stream biota.

Many alpine areas are typified by sharp gradients in environmental conditions, giving rise to ecosystem patchiness and thereby isolation (Jeník, 1997). As a result of differences in altitude and thereby climate

Chironominae Other Ephemeroptera & Plecoptera Baetidae Trichoptera Number of zoobenthic taxa and zoobenthic biomass Nemouridae Other taxa Chloroperlidae Channel stability Diamesinae Orthocladinae Simuliidae Diamesa Distance downstream from glacier margins Low Time since deglaciation Increase in water temperature Tmax= 2°C T<sub>max = 4°C</sub>

Orthocladinae

Fig. 1 Conceptual model proposed by Milner & Petts (1994) for the distribution of benthic macroinvertebrates downstream of glacial sources.

there are major changes along alpine rivers both in riparian vegetation (Lillehammer & Brittain, 1978; Petersen et al., 1995; Jeník, 1997; Zah & Uehlinger, 2001) and in instream conditions such as temperature (Burgherr & Ward, 2001; Malard et al., in press). Different water sources, for example snowmelt and glacier melt as well as the extent and nature of groundwater inputs, can also give rise to major environmental gradients over short distances which are reflected in faunal composition (Malard et al., 1999; Füreder et al., 2001). The winter period of ice and snow cover also significantly reduces the potential window for dispersal (Schütz et al., 2001). The problem of isolation also applies to several arctic and subarctic areas, both island archipelagos and nonglaciated areas between glaciers. Nevertheless, the heterogeneous nature of stream types in such catchments, particularly in floodplain areas (Gurnell et al., 1999; Smith et al., 2001), provides potential sources of colonizers to the zoobenthos. Dispersal can take place via drift of immature stages and adult flight so that the response of zoobenthic communities within the main glacial river to hydrological change can be rapid. Recent research on invertebrate colonization following glacial retreat (Milner, 1994) support this contention, although more information is required on the dispersal ability and degree of population interaction within and between catchments of this nature.

# Historical changes

The stream ecosystems of cold regions are 'young' in geological terms; they are also relatively rare today.

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During the Pleistocene, continental glaciers covered about 32% of the earth's land surface and glacial rivers would have been widespread (Embleton & King, 1975; Goudie, 1992). Today, only some 10% of the land surface remains covered by glaciers. Glacierfed rivers have been replaced by snowmelt and rainfall dominated rivers; a change reflected by adjustments of channel morphology and one that has been well-documented by palaeoenvironmental studies of river valleys throughout the temperate zone (e.g. Starkel *et al.*, 1991).

Glacial area has also varied over historic time scales. The 'Little Ice Age' between about 1550 and 1850, was the latest and most dramatic episode of Neoglaciation. The twentieth century was characterized by a general retreat of glaciers in the European Alps and in northern Europe (Fig. 2). However, the widespread retreat of glaciers in the European Alps seems to have been arrested in the 1960s and subsequently the termini of many glaciers have slowly readvanced, although overall increases in area have been small (Kasser, 1983). In northern Europe a similar retreat has been documented throughout much of the last century. However, many of the western glaciers in Norway have also begun to advance, some of them dramatically, over the past few years because of increased winter precipitation (Fig. 3, Winkler et al., 1997). In North America, there has been a general pattern of glacial retreat over the past 50 years, although in Alaska there was an increase in glacial mass in the 1970s (Mayo & March, 1990). However, this increase was not sustained (Melack et al., 1997).



**Fig. 2** The Rhone Glacier in 1825 (a) and in 1998 (b), illustrating the dramatic changes during the last two centuries. The glacier has retreated almost to the top of the steep rock face, thereby creating a glacial floodplain with another glacier-fed river, the Mutt, leading to a substantial increase in spatial and temporal habitat diversity.



Fig. 3 Patterns of glacial retreat and advance shown by selected European glaciers (Data supplied by the Norwegian Water Resources and Energy Directorate (Briksdal); National Energy Authority, Iceland (Sátujökul); Versuchsanstalt für Wasserbau Hydrologie und Glaziologie of the Swiss Federal Institute of Technology (ETH), Zurich (Mutt).

#### Anthropogenic impacts

In the future, climatic changes will influence the hydrological regime of arctic, alpine and mountain streams in a variety of ways, but the influence on glacier meltwater may be particularly significant (McGregor et al., 1995). Changes in the magnitude and variability of a range of climate determinants of glacier behaviour, as a result of an enhanced Greenhouse effect, will have important implications for the future hydrogeomorphological and thermal dynamics of glacial streams. Water resources developments, such as water diversions for hydro-electric power developments, water abstraction and land-use changes, can also have marked impacts upon these stream ecosystems, particularly on benthic communities (e.g. Boon, 1988; Brittain & Saltveit, 1989; Petts & Bickerton, 1994). Stream macroinvertebrate communities are widely used to monitor changes in water quality as a result of human impact, notably with regard to organic pollution and acidification (Rosenberg & Resh, 1993). However, knowledge of environmental variables determining macroinvertebrate distribution is a necessary prerequisite for using zoobenthos as bioindicators of environmental change in arctic and alpine running waters (Fig. 2).

Increased glacial runoff may have a significant effect on flow, temperature and sediment transport regimes. Reduced stream temperatures from increased contributions of glacial meltwater and decreased channel stability from changed runoff patterns and altered sediment loads will potentially reduce the diversity of zoobenthic communities in glacier-melt dominated rivers (McGregor *et al.*, 1995). This may cause an increase in the relative abundance of a number of key taxa, most notably *Diamesa* spp., which may thus act as a potential key indicator taxa for climate change effects in glacier-fed rivers (Melack *et al.*, 1997).

#### **Ecological processes**

The functioning of glacier-fed stream and river ecosystems in arctic and alpine regions is less well known than that of forest and lowland systems for a number of reasons. Studies have historically been largely restricted to the window of the ice-free season, because of the logistical problems of sampling in the winter. However, significant progress over the last decade has helped to redress this balance. Although these systems initially appear unstable, characterized by lower biodiversity than lowland clearwater streams, glacier-fed stream ecosystems may be unique among lotic systems in supporting predictable macroinvertebrate communities because of the overriding dominance of physical variables in determining community structure. Certain taxa groups, such as Chironomidae, may display considerable species richness and notably glacial floodplains can represent hot-spots of environmental and species diversity in view of their considerable temporal and spatial heterogeneity (Tockner et al., 1997; Ward et al., 1998). The generally low complexity and high degree of predictability of glacially influenced systems may also be an advantage, not only making them better indicators of environmental change, but also as a suitable testing ground for examining ecological concepts. The ability of the macroinvertebrates of glacier-fed streams and river ecosystems to cope with such environmental heterogeneity, through a range of life cycle and colonization strategies, indicates that their distribution patterns, in a similar manner to their lowland counterparts, are a sensitive indicator of both natural and anthropogenic changes.

# An overview of recent research summarized in this special issue

One of the recent main research thrusts has been the international co-operative project entitled, 'Arctic and Alpine Stream Ecosystem Research (AASER)' (Brittain

et al., 1998) which was the first attempt to undertake a co-ordinated study of a wide variety of glacier-fed stream ecosystems in Europe. The study was a co-operative effort between several institutions: the University of Oslo, Norway (Freshwater Ecology and Inland Fisheries Laboratory), the University of Birmingham, U.K. (School of Geography and Environmental Sciences), the University of Geneva, Switzerland (Laboratory of Ecology and Aquatic Biology), the University of Iceland, Reykjavik (Institute of Biology), the Trento Museum of Natural Sciences in Italy and the National Energy Authority of Iceland. The primary aim of AASER was to identify the primary physical and chemical variables determining the distribution of macroinvertebrates in glacially influenced catchments along a latitudinal and climatic gradient across Europe with which to test the validity of the Milner & Petts (1994) conceptual model outlined earlier in this paper. The AASER project, financed by the European Union and the Swiss Federal Office for Education and Science, covered 3 years of study from 1996 to 1999.

The AASER study sites were selected to represent a latitudinal gradient from 43°N in the French Pyrenees to 79°N on the arctic archipelago of Svalbard, and in degree of continentality from the oceanic climates of Iceland and Western Norway to the more continental climates of the Alps and Eastern Norway (Table 1, Brittain et al., 2001; Castella et al., 2001; Gislason et al., 2001; Lods-Crozet et al., 2001b; Maiolini & Lencioni, 2001; Saltveit et al., 2001; Smith et al., 2001; Snook & Milner, 2001). The study areas also represented the types of glacial situation present in Europe at present: retreating, advancing and 'stable'. Faunal diversity was also very different across sites being highest in the Alps and the Pyrenees, decreasing northwards. On Svalbard, isolated from mainland Europe, many taxa were not found, and Chironomidae dominated the freshwater fauna, both in glacial and non-glacial systems (Lods-Crozet et al., 2001a). Some systems contained lakes as well as substantial tributary influence (Brittain et al., 2001; Saltveit et al., 2001), enabling evaluation of their effect on the pattern of macroinvertebrate distribution.

Fieldwork and subsequent analyses were carried out according to a standard protocol (Brittain & Saltveit, 1997) to ensure uniformity in data collection. Within each catchment, climatic and hydrological data were collected (including temperature, solutes and suspended sediments); physical habitat characteristics were assessed (geomorphological assessment); and epilithic algae and zoobenthos were sampled. All data were entered into a single database to facilitate inter-site comparisons.

In addition to the AASER project, recent research on the zoobenthos of glacial rivers has been undertaken in the Bernia Massif of Switzerland and in the Tyrolean Alps, Austria. The Swiss site was the Val Roseg catchment, which contains a complex floodplain, providing an extensive interplay between glacial, groundwater and snowmelt channels (Malard et al., 1999). Wide-ranging studies have made a significant contribution to our understanding of both abiotic and biotic processes in glacial floodplains (e.g. Burgherr & Ward, 2001; Zah & Uehlinger, 2001). In the Tyrol, studies have been undertaken in two parallel catchments, one spring-fed (Königsbach) and one glacier-fed (Rotmoosache) (Füreder et al., 2001; Schütz et al., 2001). Both these studies and a more general study encompassing several sites throughout the Swiss Alps (Robinson et al., 2001) have included sampling during winter. Much of our knowledge of glacial rivers is based on data collected during the spring, summer and autumn. However, despite logistical problems, only year-round studies can give us the correct understanding of overall ecosystem structure and functioning and the life history patterns of the aquatic fauna.

Although recent research on glacier-fed rivers has been concentrated on European systems, studies have recently been undertaken in other areas of the world including New Zealand (Milner *et al.*, 2001) and Greenland (Friberg *et al.*, 2001). These studies widen the geographical scope of the European studies above and also provide insight into how deterministic patterns are in the downstream distribution of benthic macroinvertebrates across regions with substantially different faunas. The studies documented in this issue culminate in a synthesis paper that incorporates the development of a modified conceptual model, supported by quantitative data, that is applicable to a wide range of glacier-fed systems (Milner *et al.*, 2001).

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					Catchment					Altitudinal		
River system	Code	Region	Co-ordinates	Max. altitude of catchment (m a.s.l.)	area at downstream reach (km <sup>2</sup> )	Glacier area (km <sup>2</sup> )	Precipiation (mm), 1996/97	Number of study reaches	Distance of reaches from glacier (m)	range of river studied (m a.s.l.)	Discharge range (m <sup>3</sup> s <sup>-1</sup> )	Treeline altitude (m a.s.l.)
Taillon	TAI	Pyrenees	43°06'N, 0°01'W	2975	6.4	0.2	+	4	50-1500	1870–2500	0.3-0.4	1600
Conca	CON	S.E. Alps	46°06'N, 10°36'F	3463	4.2	0.2	1497	~	350-4600	1300–2830	0.3–2.1	2000
Mutt	MUT	N. Alps	46°33'N,	3099	~	9.0	1595	Ŋ	5-3600	1800–2600	0.6–1.5	2000
Dalelva	BRI	W. Norway	0 24 E 61°40'N, 650'E	1915	25.6	22*	1270	Ŋ	100-7100	10–340	0.5–35	600
Leirungsåi	LEI	Central	61°24'N,	2159	400	1.2	с. 800	Ч	200-24600	970-1550	0.4–6.2	1050
W-Jokülsá	0ĺM	Inorway Central Icolond	8*41 E 64°50'N, 18°45'M	1800	840	68	662	11	5-45000	160-860	25–31	No traact
Bayelva	ВАҮ	Spitsbergen, Svalbard	10 ±0 W 78°55′N, 11°50′E	742	30.9	12	483	4	300-2900	5-50	0.01–15.6	No trees
*Part of a lat tc. 300 mm 1 ‡The natural	ge plate ecorded treeline,	au glacier with during July an , without anthr	an area of 487 l id August. opogenic impact	cm <sup>2</sup> . ts, would be 300-	400 m a.s.l.							

Table 1 Location and characteristics of the glacier-fed rivers investigated in the AASER project

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